

Package ‘bayesianETAS’

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Title Bayesian Estimation of the ETAS Model for Earthquake Occurrences

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Description The Epidemic Type Aftershock Sequence (ETAS) model is one of the best-performing methods for modeling and forecasting earthquake occurrences. This package implements Bayesian estimation routines to draw samples from the full posterior distribution of the model parameters, given an earthquake catalog. The paper on which this package is based is Gordon J. Ross - Bayesian Estimation of the ETAS Model for Earthquake Occurrences (2016), available from the below URL.

URL <http://www.gordonjross.co.uk/bayesianetas.pdf>

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 bayesianETAS

Bayesian estimation of the ETAS model for earthquake occurrences

Description

Bayesian estimation of the ETAS model for earthquake occurrences

Author(s)

Gordon J Ross <gordon@gordonjross.co.uk>

References

Gordon J. Ross - Bayesian Estimation of the ETAS Model for Earthquake Occurrences (2016), available from <http://www.gordonjross.co.uk/bayesianetas.pdf>

maxLikelihoodETAS

Estimate the parameters of the ETAS model using maximum likelihood.

Description

The Epidemic Type Aftershock Sequence (ETAS) model is widely used to quantify the degree of seismic activity in a geographical region, and to forecast the occurrence of future mainshocks and aftershocks (Ross 2016). The temporal ETAS model is a point process where the probability of an earthquake occurring at time t depends on the previous seismicity H_t , and is defined by the conditional intensity function:

$$\lambda(t|H_t) = \mu + \sum_{t[i]<t} \kappa(m[i]|K, \alpha)h(t[i]|c, p)$$

where

$$\kappa(m_i|K, \alpha) = Ke^{\alpha(m_i - M_0)}$$

and

$$h(t_i|c, p) = \frac{(p-1)c^{p-1}}{(t-t_i+c)^p}$$

where the summation is over all previous earthquakes that occurred in the region, with the i 'th such earthquake occurring at time t_i and having magnitude m_i . The quantity M_0 denotes the magnitude of completeness of the catalog, so that $m_i \geq M_0$ for all i . The temporal ETAS model has 5 parameters: μ controls the background rate of seismicity, K and α determine the productivity (average number of aftershocks) of an earthquake with magnitude m , and c and p are the parameters of the Modified Omori Law (which has here been normalized to integrate to 1) and represent the speed at which the aftershock rate decays over time. Each earthquake is assumed to have a magnitude which is an independent draw from the Gutenberg-Richter law $p(m_i) = \beta e^{\beta(m_i - M_0)}$.

This function estimates the parameters of the ETAS model using maximum likelihood

Usage

```
maxLikelihoodETAS(ts, magnitudes, M0, T, initval = NA, displayOutput = TRUE)
```

Arguments

ts	Vector containing the earthquake times
magnitudes	Vector containing the earthquake magnitudes
M0	Magnitude of completeness.
T	Length of the time window [0,T] the catalog was observed over. If not specified, will be taken as the time of the last earthquake.
initval	Initial value at which to start the estimation. A vector, with elements (mu, K, alpha, c, p)
displayOutput	If TRUE then prints the out the likelihood during model fitting.

Value

A list consisting of

params	A vector containing the estimated parameters, in the order (mu,K,alpha,c,p,beta)
loglik	The corresponding loglikelihood

Author(s)

Gordon J Ross

References

Gordon J. Ross - Bayesian Estimation of the ETAS Model for Earthquake Occurrences (2016), available from <http://www.gordonjross.co.uk/bayesianetas.pdf>

Examples

```
## Not run:  
beta <- 2.4; M0 <- 3; T <- 500  
catalog <- simulateETAS(0.2, 0.2, 1.5, 0.5, 2, beta, M0, T)  
maxLikelihoodETAS(catalog$ts, catalog$magnitudes, M0, 500)  
  
## End(Not run)
```

sampleETASposterior *Draws samples from the posterior distribution of the ETAS model*

Description

This function implements the latent variable MCMC scheme from (Ross 2016) which draws samples from the Bayesian posterior distribution of the Epidemic Type Aftershock Sequence (ETAS) model.

The ETAS model is widely used to quantify the degree of seismic activity in a geographical region, and to forecast the occurrence of future mainshocks and aftershocks (Ross 2016). The temporal ETAS model is a point process where the probability of an earthquake occurring at time t depends on the previous seismicity H_t , and is defined by the conditional intensity function:

$$\lambda(t|H_t) = \mu + \sum_{t[i]<t} \kappa(m[i]|K, \alpha)h(t[i]|c, p)$$

where

$$\kappa(m_i|K, \alpha) = Ke^{\alpha(m_i - M_0)}$$

and

$$h(t_i|c, p) = \frac{(p-1)c^{p-1}}{(t-t_i+c)^p}$$

where the summation is over all previous earthquakes that occurred in the region, with the i 'th such earthquake occurring at time t_i and having magnitude m_i . The quantity M_0 denotes the magnitude of completeness of the catalog, so that $m_i \geq M_0$ for all i . The temporal ETAS model has 5 parameters: μ controls the background rate of seismicity, K and α determine the productivity (average number of aftershocks) of an earthquake with magnitude m , and c and p are the parameters of the Modified Omori Law (which has here been normalized to integrate to 1) and represent the speed at which the aftershock rate decays over time. Each earthquake is assumed to have a magnitude which is an independent draw from the Gutenberg-Richter law $p(m_i) = \beta e^{\beta(m_i - M_0)}$.

Usage

```
sampleETASposterior(ts, magnitudes, M0, T = NA, initval = NA,
  approx = FALSE, sims = 5000, burnin = 500)
```

Arguments

ts	Vector containing the earthquake times
magnitudes	Vector containing the earthquake magnitudes
M0	Magnitude of completeness.
T	Length of the time window [0,T] the catalog was observed over. If not specified, will be taken as the time of the last earthquake.
initval	Initial value at which to start the estimation. If specified, should be a vector, with elements (mu, K, alpha, c, p). If unspecified, the sampler will be initialized at the maximum likelihood estimate of the model parameters

approx	If TRUE then will approximate the true posterior using the infinite time approximation discussed in (Ross 2016)
sims	Number of posterior samples to draw
burnin	Number of burnin samples

Value

A matrix containing the posterior samples. Each row is a single sample, and the columns correspond to (mu, K, alpha, c, p)

Author(s)

Gordon J Ross

References

Gordon J. Ross - Bayesian Estimation of the ETAS Model for Earthquake Occurrences (2016), available from <http://www.gordonjross.co.uk/bayesianetas.pdf>

Examples

```
## Not run:
beta <- 2.4; M0 <- 3; T <- 500
catalog <- simulateETAS(0.2, 0.2, 1.5, 0.5, 2, beta, M0, T)
sampleETASposterior(catalog$ts, catalog$magnitudes, M0, T, sims=5000)

## End(Not run)
```

simulateETAS

Simulates synthetic data from the ETAS model

Description

This function simulates sample data from the ETAS model over a particular interval $[0, T]$. The Epidemic Type Aftershock Sequence (ETAS) model is widely used to quantify the degree of seismic activity in a geographical region, and to forecast the occurrence of future mainshocks and aftershocks (Ross 2016). The temporal ETAS model is a point process where the probability of an earthquake occurring at time t depends on the previous seismicity H_t , and is defined by the conditional intensity function:

$$\lambda(t|H_t) = \mu + \sum_{t[i] < t} \kappa(m[i]|K, \alpha) h(t[i]|c, p)$$

where

$$\kappa(m_i|K, \alpha) = K e^{\alpha(m_i - M_0)}$$

and

$$h(t_i|c, p) = \frac{(p-1)c^{p-1}}{(t-t_i+c)^p}$$

where the summation is over all previous earthquakes that occurred in the region, with the i 'th such earthquake occurring at time t_i and having magnitude m_i . The quantity M_0 denotes the magnitude of completeness of the catalog, so that $m_i \geq M_0$ for all i . The temporal ETAS model has 5 parameters: μ controls the background rate of seismicity, K and α determine the productivity (average number of aftershocks) of an earthquake with magnitude m , and c and p are the parameters of the Modified Omori Law (which has here been normalized to integrate to 1) and represent the speed at which the aftershock rate decays over time. Each earthquake is assumed to have a magnitude which is an independent draw from the Gutenberg-Richter law $p(m_i) = \beta e^{\beta(m_i - M_0)}$.

This function simulates sample data from the ETAS model over a particular interval $[0, T]$.

Usage

```
simulateETAS(mu, K, alpha, c, p, beta, M0, T, displayOutput = TRUE)
```

Arguments

mu	Parameter of the ETAS model as described above.
K	Parameter of the ETAS model as described above.
alpha	Parameter of the ETAS model as described above.
c	Parameter of the ETAS model as described above.
p	Parameter of the ETAS model as described above.
beta	Parameter of the Gutenberg-Richter law used to generate earthquake magnitudes.
M0	Magnitude of completeness.
T	Length of the time window $[0, T]$ to simulate the catalog over.
displayOutput	If TRUE then prints the number of earthquakes simulated so far.

Value

A list consisting of

ts	The simulated earthquake times
magnitudes	The simulated earthquake magnitudes
branching	The simulated branching structure, where branching[i] is the index of the earthquake that triggered earthquake i , or 0 if earthquake i is a background event

Author(s)

Gordon J Ross

References

Gordon J. Ross - Bayesian Estimation of the ETAS Model for Earthquake Occurrences (2016), available from <http://www.gordonjross.co.uk/bayesianetas.pdf>

Examples

```
## Not run:
beta <- 2.4; M0 <- 3
simulateETAS(0.2, 0.2, 1.5, 0.5, 2, beta, M0, T=500, displayOutput=FALSE)

## End(Not run)
```

simulateNHPP	<i>Simulates event times from an inhomogenous Poisson process on [0,T]</i>
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Description

Simulates event times from an inhomogenous Poisson process on [0,T]

Usage

```
simulateNHPP(targetfn, maxintensity, T = Inf)
```

Arguments

targetfn	A first order function defining the process intensity
maxintensity	The maximum values of targetfn
T	Length of the interval [0,T] on which to simulate the process

Value

The simulated event times

Author(s)

Gordon J Ross

Examples

```
simulateNHPP(function(x) {sin(x)+1}, 2, 100)
simulateNHPP(function(x) {x^2}, 100, 10)
```

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